



AN IONOSPHERIC NUMERICAL MODEL AND SOME RESULTS FOR THE ELECTRON DENSITY STRUCTURE BELOW THE F₂ PEAK

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ABSTRACT

An aeronomical theoretical model fitted to the observed peak data reproduces many observed features including the appearance of an F₁-ledge.

SUMMARY OF THEORY AND RESULTS

A numerical model for the mid-latitude ionospheric profile below the F₂-peak is constructed /1,2/, based on solving a set of equations for the multi-composition system involving O⁺, NO⁺, O₂⁺ and N₂⁺, as well as electrons. Dynamic processes of diffusion and wind, and also photo-chemical processes, with 21 reactions for stable as well as meta-stable ions of atomic oxygen are taken into account. The EUV91 model /3/ is used to specify the solar irradiance. Neutral atmospheric density and temperature are provided by the MSIS86 model /4/. The topside O⁺ density has been taken from empirical models, e. g., IRI90 /5/, and the neutral meridional wind from HWM90 /6/. Moreover, in order to gain a further realistic ionospheric profile, the model uses the observed foF₂ and the derived hmF₂ (from M3000F₂) as input so that the key point on the simulated vertical distribution curve of electron density, namely the F₂-layer peak is the observed value. The model values for the topside density and the meridional wind can be adjusted and even replaced. Thus the resultant profile may be used readily for comparison with the observed ionospheric profile without normalization, and provide useful information for the ionogram inversion when valleys exist and for the empirical modelling of the profile shape like IRI. With the help of the servo model /7/, vertical drifts can be derived from ionosonde data while assuming an equilibrium state of the ionosphere /8/. Using this ionosonde data based drift helps to obtain a more reliable ionospheric profile than using an empirical horizontal wind model like HWM90.

Figure 1 (a) shows the diurnal variation of the electron density profile for Wakkanai (45.6°N, 141.7°E), June 1982. The F₁-ledge becomes evident with the foF₁ increase and the hmF₁ decrease toward noon. The seasonal change of the profile, as shown in Figure 2(b) where the observed foF₂ and hmF₂ are introduced (i. e. the topside density is modified according to foF₂, and derived vertical drifts are used instead of HWM90), reveals that the F₁-ledge disappears in winter and is much more pronounced in summer under this high solar activity condition. It has been found also the ledge is fully developed near noon in summer for low solar activity.

Peak heights hmF₂ are compared for two set-ups: (1) when HWM90 is used; (2) when the derived vertical drift is used. It appears (see, for example, Figure 2) that there exists much better agreement between the simulated hmF₂ and the observation when derived drifts are used than when HWM90 is used. This implies that the servo model based drift contains very practical information on the ionospheric dynamics so that it can be used to reproduce the hmF₂.

Vertical distributions of the atomic ion O^+ and molecular ions share a common density at the ion transition level situated by day around 180 to 190 km (Figure 3). The present model gives an obvious diurnal change, while the seasonal and the solar activity changes are not so strong as expected.

The model gives also results for the E-F valley size, the width and the depth of the valley. It is found that the width is by day generally around 12 km for low solar activity and slightly wider for high solar activity. The depth, defined as the ratio between the plasma frequency at the bottom of the valley and the critical frequency of the E-layer, is about 0.985 to 0.990 with a larger value at high solar activity.

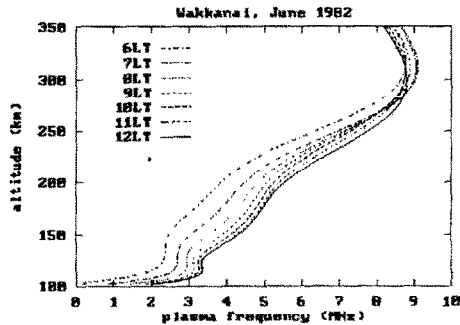


Fig. 1 (a) Diurnal variation of the electron density profile

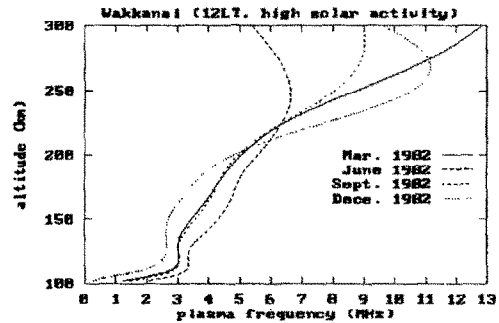


Fig. 1(b) Seasonal variation of computed profile (foF2 and hmF2 are used)

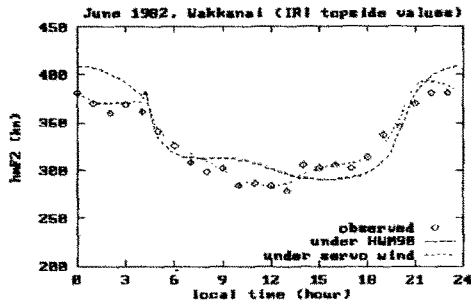


Fig. 2 hmF2 diurnal variation observed vs. computed using HWM90 wind and servo model based drifts

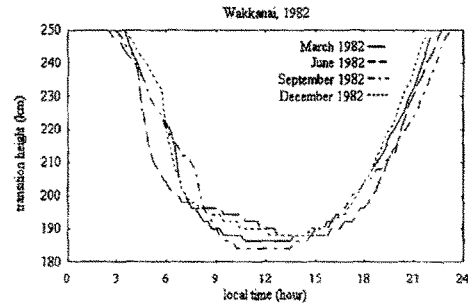


Fig. 3 Variation of the ion transition height with local time for Wakkanaï, 1982

REFERENCES

1. Zhang, S.-R., X.-Y. Huang, Y.-Z. Su and S. M. Radicella, *Annali di Geofisica*, 36, 105-110 (1993)
2. Zhang, S.-R. and X.-Y. Huang, *ACTA Chinese J. Space Sci.*, 14 (1994) (in press)
3. Tobiska, K., *J. Atmos. Terr. Phys.*, 55, 1637-1659 (1993)
4. Hedin, A.E., *J. Geophys. Res.*, 92, 4649-4662 (1987)
5. Bilitza, D., *NSSDC/WDC-A for Rockets and Satellites 90-22*, (1990)
6. Hedin, A. E., M. A. Biondi, R. G. Burnside, G. Hernandez, R. M. Johnson, T. L. Kileen, C. Mazaudier, J. W. Meriwether, J. E. Salah, R. J. Sica, R. W. Smith, N. W. Spencer, V. B. Wickwar and T. S. Virdi, *J. Geophys. Res.*, 96, 7657-7688 (1991)
7. Rishbeth, H., S. Ganguly and J. C. G. Walker, *J. Atmos. Terr. Phys.*, 40, 767-784 (1978)
8. Titheridge, J. E., *J. Atmos. Terr. Phys.*, 55, 1637-1659 (1993)